The role of nitrate radicals (NO₃) in aerosol life cycle: Secondary organic aerosol formation and aging of atmospheric organic aerosols

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Nitrate Radical Oxidation of BVOCs

- Nitrate radicals (NO₃) are a major night time oxidant
 - Reacts with approximately 20% of all biogenic volatile organic carbons (Pye et al., 2010)
- Monoterpene+NO₃ reaction has high organic nitrate and SOA yields (e.g., Griffin et al., 1999; Fry et al., 2009)
- NO₃ is created by the reaction of anthropogenic NO₂ with O₃

$$NO_2 + O_3 \rightarrow NO_3 + O_2$$



Nitrate radical oxidation of BVOCs represent a direct way for linking anthropogenic emissions and biogenic SOA formation

Synergetic Approach of Studying Organic Aerosols

Research Goal: determine the extent to which NO_3 radicals oxidation of BVOC affect organic aerosol loading and composition over its atmospheric lifetime



Field Measurements: SOAS and SCAPE

- Southern Oxidant and Aerosol Study (SOAS)
 - 2013 June July
 - Centreville (rural Alabama)

- Southeastern Center of Air Pollution and Epidemiology study (SCAPE)
 - 2012 May 2013 Feb
 - Greater Atlanta Area (urban and rural)



OA Source Apportionment in the SE US



- OA sources vary spatially and seasonally
- Less-Oxidized Oxygenated Organic Aerosol (LO-OOA) is an important factor across all seasons and sites
 - LO-OOA can account for 19-34% of total OA in SE US

Xu et al., PNAS, 2015

LO-OOA Source: local production or long-range transport?



- Diurnal trends are similar for all sites and all seasons (diurnal trends are normalized by the highest LO-OOA concentration of each dataset).
- A diurnal maximum at night and a minimum in the afternoon.

Xu et al., PNAS, 2015

LO-OOA Nighttime Increase

-- boundary layer height change or production?



Nighttime increase still exists after adjusting by the boundary layer height → nighttime aerosol production

Xu et al., PNAS, 2015

LO-OOA at SOAS



- LO-OOA peaks at night and has same diurnal as monoterpenes.
- LO-OOA is identified in all seasons → monoterpene seasonal variation

- LO-OOA is strongly correlated with "nitrate groups (-ONO2) in organic nitrates"
- Estimated based on AMS-IC method

Contribution of monoterpenes + NO₃• chemistry to LO-OOA

Estimation of Particulate Organic Nitrates

Three independent methods to estimate particulate organic nitrates

1. AMS-IC method: AMS total nitrate - PILS inorganic nitrate



3. PMF method: include NO⁺ and NO₂⁺ in PMF analysis

Xu et al., ACP, 2015

NO3_{org} and NO3_{inorg} in the SE US

- Concentration of "nitrate groups" (-ONO2)
- Organic origin: similar amount year round, is ~ 0.2 ug/m³
- Inorganic origin: higher in winter months, ~ 0.8 1.4 ug/m³



Xu et al., ACP, 2015

Ubiquitous Presence of Particulate Organic Nitrate



Organic origin: 63-100% of total "nitrate groups" conc. in summer

Organic nitrates are 5-12% of total OA in summer (assume MW = 200 -300 g/mole)

Xu et al., ACP, 2015

Fundamental Lab Studies: Monoterpenes + NO₃

- 1. Laboratory studies of SOA formation from monoterpenes + NO₃ SOA (SOA yields, formation mechanisms, organic nitrates)
 - Effect of RH
 - Seed Acidity (highly acidic seed, Guo et al., 2015)
 - Peroxy radical fate: Mostly likely "Low-NOx" chemistry (RO₂+HO₂)
 - Loadings ~ 10 µg/m³
 - α -pinene, β -pinene, limonene

2. Changes in SOA and organic nitrates with continued processing

- Hydrolysis
- Dilution
- Temperature Change
- Dark/photochemical aging

Fate of Peroxy Radicals (RO₂)



Figure 1. General schematic of gas-phase peroxy radical chemistry in SOA formation.

Continued Processing of OA and ON

- Different Perturbations
- Changes with hydrolysis, atmospheric dilution, temperature change, photochemical (OH) and dark (NO₃) aging,



Georgia Tech Environmental Chamber Facility

Dual chamber facility, 300 lights, temperature range 4- 40 °C.



Georgia Tech Environmental Chamber Facility



Gas-phase measurements

- Q-CIMS
- HR-ToF-CIMS
- GC-FID
- O₃ monitor
- Chemilumnescence NO/NO₂/NO_x analyzer
- CAPS NO₂ monitor

Particle-phase measurements

- HR-ToF-AMS
- FIGAERO-HR-ToF-CIMS
- SMPS
- CPC
- Offline filter characterization

SOA Formation from α -pinene + NO₃ and β -pinene + NO₃



- Temperature = 25 °C
- RH = 50 54%
- Highly acidic MgSO₄/H₂SO₄ seed
- RO₂+HO₂ and RO₂+NO₃ pathways

Nah et al., ES&T, 2016

β -pinene+NO₃: SOA Yields



- Humidity and RO₂ fate does not have a strong effect on aerosol mass yield
- Aerosol mass yields: 27.0-104.1% for mass loadings ranging from 5.1-216.1 $\mu g/m^3$
- β -pinene+NO₃ can potentially contribute to a large fraction of ambient aerosol

β -pinene+NO₃ : Gas-phase Products



- Aerosol and gas phase species appear almost immediately
- Organic nitrate species identified by Q-CIMS
- Fast reaction with immediate condensation of products



β -pinene+NO₃: Aerosol Composition and Dark Aging



- Large fraction of nitrate species at NO⁺ (m/z 30) and NO₂⁺ (m/z 46), 11%
- NO^+ / NO_2^+ ratio = 4.8 10.2
- Relatively large signal at m/z 67 ($C_5H_7^+$) and m/z 91 ($C_7H_7^+$)
- Increase in O/C with dark aging: 18% (dry), 6% (humid)

Relevance of β -pinene+NO₃ to SOAS

Spectrum of LO-OOA at SOAS has similar features to laboratory β -pinene+NO₃ SOA at m/z > 60



Highly-Oxygenated ON in Chamber Studies (β-pinene + NO₃)



- Highly-oxygenated ON observed in FIGAERO-CIMS with 4 – 9 oxygen atoms
- Rapid formation, likely occurs through auto-oxidation (Crounse et al., 2013; Ehn et al., 2014)

Nah et al., ES&T, 2016

ON Observed with FIGAERO-HR-ToF-CIMS

361.937836

C-HaNO

β -pinene + NO₃

 α -pinene + NO₃

				-
Molecule-iodide adduct formula	Exact m/z		Molecule-iodide adduct formula	Exact m/z
C10H17NOd	342.020782		C10H12NO2	356.000031
C10H15NO5I	356.000031		C10H15NO5	371.994965
C ₁₀ H ₁₇ NO ₅ I	358.015686		C10H16N2O7	403.000763
C10H19NO5I	360.031311		C10H12NO2	419,979706
C ₁₀ H ₁₂ NO ₆ I	369.97934		C ₄ H ₁ NO ₄	357.979309
C10H15NO6I	371.994965		Cationod	557.979509
C10H17NO6I	374.01062			
C10H12NOgI	376.026306			
C10H12NO7I	385.974213			
C10H15NO7I	387.989868			
C10H17NO7I	390.005524	· 1'	1 ON	
C10H19NO-I	392.02121	• 4.	1 ON observed fo	$^{\circ}$ B-pinene+NO ₂
C ₁₀ H13NO ₂ I	401.969147			II 5
C ₁₀ H ₁₅ NO ₅ I	403.984772	10	2 abconvod at CO	AS, Lee et al., 2016
C ₁₀ H ₁₇ NO ₅ I	406.000427	(3	z observed at 50	AS, LEE EL dI., ZUID
C10H12NO2	408.016052	•		
C10H12NO9I	417.96405			
C10H15NO9I	<mark>419.979675</mark>		ON for a ninonal	
C10H17NO9I	421.995361	• 5	ON for α -pinene+	
C ₂ H ₁₃ NO ₂ I	357.979309		•	J
C ₂ H ₁₂ NO ₂ I	359.994965		observed at SOA	S, Lee et al., 2016)
C ₉ H ₁₂ NO ₇ I [*]	<mark>373.974182</mark>	(4	· Ubserveu at SUA	$\mathbf{J}, \mathbf{L} \in \mathbf{C} \subset \mathbf{I}$
C ₉ H ₁₂ NO ₇ I	375.989868			
C ₂ H ₁₇ NO ₇ I	378.005524			
C ₂ H ₁₂ NO ₂ I	389.969147			
C ₂ H ₁₂ NO ₂ I	391.984833			
C ₂ H ₁₂ NO ₂ I	405.96405			
CoH12NOp1	407.979706		chemistry forms	atmospherically
C ₂ H ₁₁ NO ₂ I	343.963654		chemistry forms	aunosphericany
C _s H ₁₃ INO _s I	345.979309			
C ₂ H ₁₁ NO ₂ T	<mark>359.958588</mark>	rele	vant highly oxyge	nated ON
C ₂ H ₁₂ NO ₂ T	361.974213			
C ₂ H ₁₁ NO ₂ T	375.953491			
C ₁ H ₁₁ NO ₁ I	377.969147			
C ₂ H ₁₁ NO ₂ I	391.948395			
C ₇ H ₆ NO ₅ I	313.953094			
C ₇ H ₆ NO ₆ I	329.947998			
C7H11NO6I	331.963654			
C-HaNO-I	<u>345.942932</u>			Nobotal ECOT 2016
C ₇ H ₁₁ NO ₇ I	<mark>347.958588</mark>			Nah et al., ES&T, 2016

Fate of Atmospheric ON



 Aerosol ON that is not removed by deposition can either form nitric acid through hydrolysis or be released as NO_x through photolysis or OH oxidation

1. β -pinene+NO₃: Organic Nitrate Hydrolysis



- Likely that majority (~90%) of ON is primary (which do not hydrolyze at appreciable rates)
- ~ 10% of ON hydrolyzes with a lifetime of 3-4.5 hr
 - Much higher than primary/tertiary ON ratio predicted by SOA formed from photooxidation under high NOx conditions (Browne et al., 2013)
- What happens to ON that do not hydrolyze? (do they get photolyzed /OH reaction?)

2. Photochemical Aging of Nighttime Aerosol

 β -pinene + NO₃

 α -pinene + NO₃



- Photochemically aging, OH + hບ
- β-pinene+NO₃ SOA: photochemical aging has little effect
 - α-pinene+NO₃:
 a large fraction of reactive
 nitrogen is released from the
 α-pinene SOA back to the gas
 phase during photooxidation.

Nah et al., ES&T, 2016

Conclusions

- LO-OOA (19-34%) : likely originates from monoterpenes, and its formation could be controlled by nighttime NO₃ chemistry (NO_x effect)
- ON contributes up to 12% of ambient OA in SE US in summer
- β -pinene+NO₃ reaction
 - FIGAERO-HR-ToF-CIMS: Formation of highly oxygenated ON species in both gas- and particlephase, many of which observed at SOAS
 - This reaction likely contributes substantially to ambient LO-OOA and ON
 - Peroxy radical fate and RH does not have a strong effect on SOA yield
- Hydrolysis: ~90% of the ON formed from the β-pinene+NO₃ reaction are primary nitrates, do not hydrolyze at appreciable rates
- Atmospheric perturbations
 - Particulate ON from β-pinene+NO₃: NOx sinks
 - Particulate ON from α-pinene+NO₃: photochemical oxidation may be an important atmospheric NOx source in the day

Laboratory studies should be an integrated part of field studies

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